

01.17 Geological Outline (2013 Edition)

Overview

Geological maps are a synthesis of general geo-scientific information on the present situation, and involve an evaluation of all drilling results, ascertainments, the morphology and of the state of knowledge to date. In addition, there are maps on particular specific areas derived from or based upon geological maps, such as [Engineering-geological maps](#), [Groundwater isomaps](#), or [Maps of the geothermic potential](#).

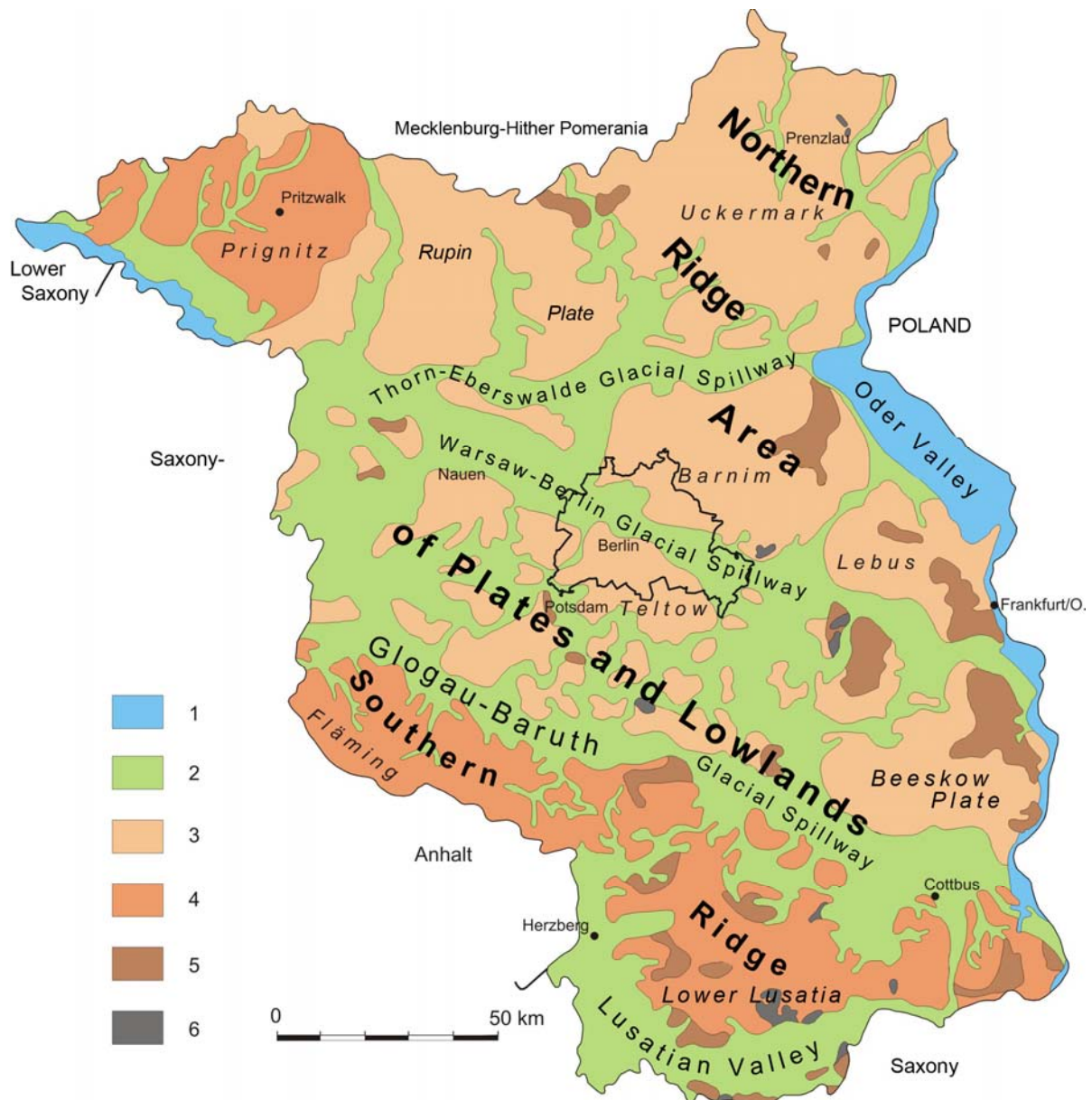
Since geological maps can generally depict only those geological units which are present at the surface, they must be supplemented by information on the sequence of layers, the three-dimensional structure of a profile section, and an explanatory text such as the present one.

Due to the large degree of compilation of various geological units for the sake of better visibility, this geological overview map has been called a **Geological Outline**.

The morphology of the Berlin-Brandenburg landscape and its development

The **basic pattern of the surface formation of Berlin and Brandenburg** is essentially determined by three major relief units which extend through the area from southeast to northwest, and which owe their formation and emergence to the processes which occurred during the Quaternary Ice Age (Fig. 1). They are:

- The Southern Ridge, including the Fläming region and the Lower Lusatian boundary wall, which to the south borders on the Lusatian Valley, itself part of the Breslau-Magdeburg Glacial Spillway
- The broad, but very heterogeneous intermediate Plates and Lowlands Area, with a large number of greater and lesser, variegated plateaus, interspersed by three large glacial spillways, the Glogau-Baruth, the Warsaw-Berlin and the Thorn-Eberswalde Spillways, and
- The Northern or Baltic Ridge, which, in Brandenburg, includes the Uckermark district.



1 – River floodplains; 2 – Glacial spillways incl. meltwater valleys, low fluvial terrace, peri-glacial basin in the old moraine area; 3 – 4 Ground moraine, melt-water formations, end moraines; 3 – Young drift plateau; 4 – Old drift plateau; 5 – Areas of old deformed deposits; 6 - Pre-quaternary

Fig. 1: Geological-morphological structure of Brandenburg and Berlin (Source: Sonntag 1995)

The emergence and formation of the natural landscape in Berlin-Brandenburg are closely connected to the most recent era of the earth's development, the **Quaternary**, generally also known as the Ice Age. The Quaternary is a segment of the earth's history during which global climatic fluctuations led to repeated shifts between glaciations and warmer interglacials in relatively short geological time periods, with very considerable impact on the fauna and the flora. During the recent phase of the Quaternary, a large-scale expansion of huge inland ice covered the northern part of central Europe. With annual mean temperatures of below 0°C, conditions during the Ice Age at our latitudes were similar to those that we would encounter today in the ice-covered regions of Greenland or Spitzbergen.

For more than 100 years, the term **Glacial Series** has been used to designate the characteristic horizontal/spatial landscape sequence and the morphological phenomenology typical of the processes during the era of the advance and the later disintegration and melting of the inland ice masses (Fig. 2).

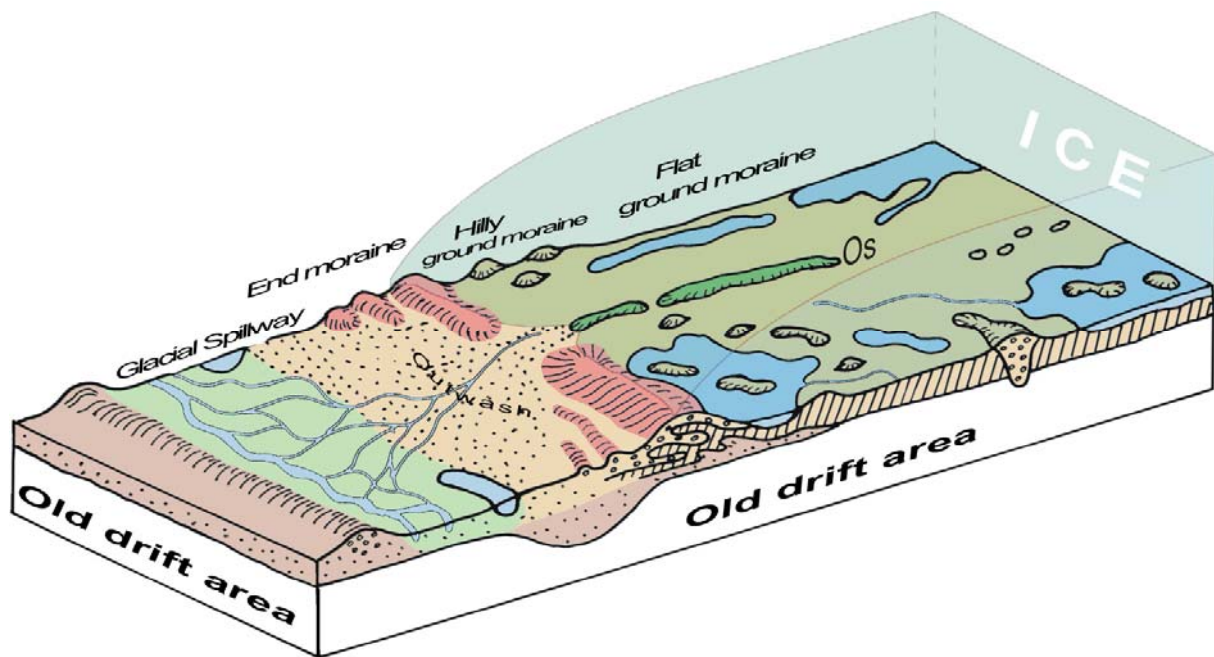


Fig. 2: The Glacial Series – a schematic image compiled into a block representation, from various authors, by Sonntag (2005)

The main geo-morphological elements of the Glacial Series are:

- The **end moraine**, formed by sediment accumulation at the edge of the ice due to buckling, or as a result of a more or less stationary glacial edge
- The **ground moraine** deposited by the ice in the rear area
- The **outwash**, deposited over a wide area in front of the glacier (the **outwash plain**) by meltwater, and finally
- The **glacial spillway**, which brings together the outflowing meltwater and passes it off toward the northwest.

The glacial series is supplemented by a number of generally less frequently encountered geomorphological formations such as **eskers**, **kames** and **drumlins**, and also **tongue bassins** and **kettle-holes**, but especially by the channels and lakes which mark the landscape (cf. MARCINEK & NITZ 1973; also LIEDTKE 1975).

Geological overview of Tertiary and Quaternary formations

The earth's surface and the near-surface segment in the area of the city of Berlin were formed by deposits during the Tertiary, the Pleistocene and the Holocene, with the tertiary and quaternary sediments having a certainly decisive significance for the life of the city. For instance, all the water from the public water supply system comes largely from the Quaternary and to some extent from the tertiary groundwater aquifers. Some 90% of the groundwater used in Berlin is withdrawn within the city limits.

The Tertiary

The Tertiary, also known as the Lignite Era, began 65 million years ago, at the end of the Cretaceous period and ended at the beginning of the Quaternary, some 2.6 million years ago.

Sediments from the Tertiary are found in Berlin in a discordant pattern, with a time gap, above older strata of material from the Cretaceous, Jurassic or Keupers eras, and constitute a virtually complete underlying, older level beneath the quaternary strata (Figure 3). In areas of deeper quaternary erosion, such as e.g. in sections of quaternary removal zones (tunnel valleys), tertiary deposits are entirely missing. Tertiary sediments constitute the surface layer at only one site in Berlin, in Lübars.

The clayey/silty rupelton, which is some 80 m thick, has a particular position within the tertiary layer. Due to its very broad dissemination, it is not only an important geological guide horizon within the covering hills from the Tertiary and Quaternary, it is also of considerable hydrogeological significance,

since it constitutes a **barrier** between the **saltwater** in the underlying stratum and the **freshwater** in the younger, covering strata. In areas in which rupelton strata are partially or completely absent due to quaternary erosion, dissemination pathways, in some cases extending to the surface, were created for the rise of mineralized deep water, and for its spread into the freshwater strata (Figure 3).

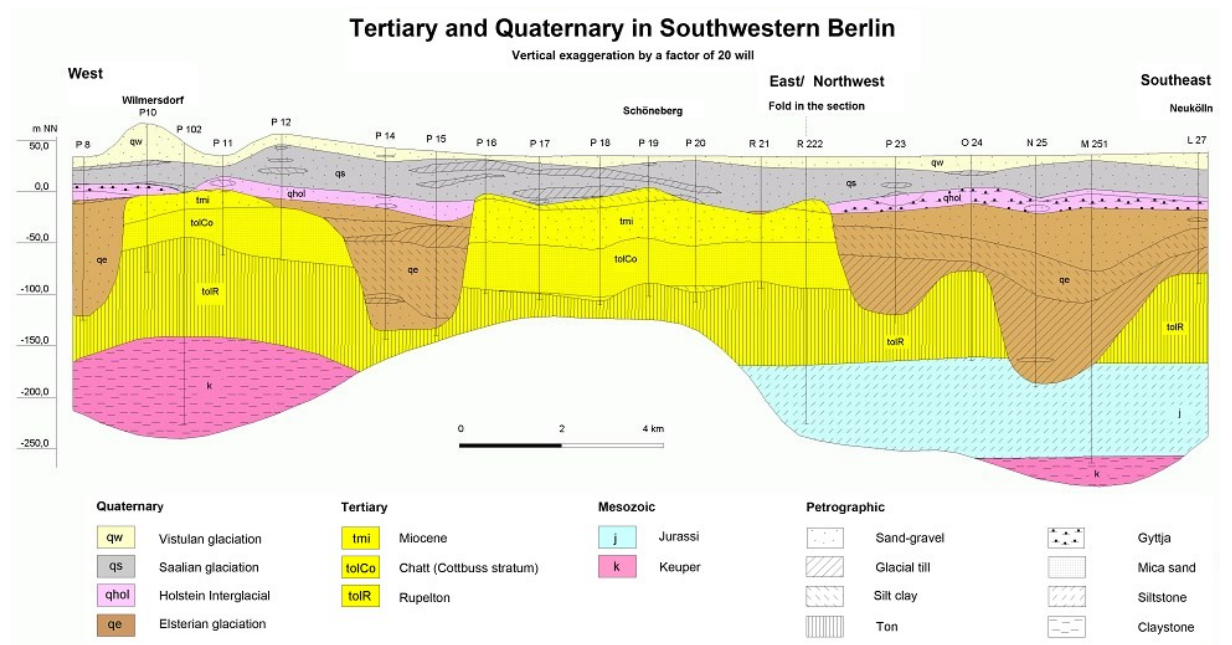


Fig. 3: Simplified geological profile through the tertiary and quaternary strata, extending from Wilmersdorf through Schöneberg to Neukölln

The Quaternary

The Quaternary began with a global climate change, the Pleistocene Ice Age, some 2.6 million years ago, and ended with the Holocene at the end of the Ice Age. In Berlin, Pleistocene deposits from the three Nordic glaciations, the Elsterian, the Saalian and the Weichselian Glaciations, occurred. These deposits consist of meltwater sands and gravels of Nordic origin, laminated clays and silts, and the glacial till of the ground moraines.

In addition, there are sediments from the Holstein and Eemian interglacials, with gyttjas, silts, clays and peats, and also river sands and gravels provided from areas further south (Table 1).

Tab. 1: Schematic structure of the Quaternary, by Stackebrandt & Manhenke (2010)			
Series	Stage (north German system)	Typical deposits	Began years before present
Holocene		Peats, gyttjas, dune and river sands, bog-associated sands	12,000
Pleistocene	Weichselian Glaciation	Meltwater sands, boulder clay, glacial till	115,000
	Eemian Interglacial	Peats, gyttjas, bog-associated sands	127,000
	Saalian Glaciation	Meltwater sands, boulder clay, glacial till, basin clays and silts	304,000
	Holstein Interglacial	Peats, gyttjas, clays and silts with snails, river sands and gravels	320,000
	Elsterian Glaciation	Meltwater sands, boulder clay, glacial till, basin clays and silts	400,000

	Older glaciations and interglacials	None observed in Berlin	2,600,000
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Tab 1: Schematic structure of the Quaternary, by Stackebrandt & Manhenke (2010)

The glaciers, particularly those of the Elsterian glaciation, eroded channels in the existing tertiary surface, in some cases very deep, which were then filled with glacial rock material. In some cases, the tertiary layer, particularly the rupelton clay, was completely eroded away, so that the protective barrier between the freshwater and saltwater strata was destroyed.

The thickness of the Pleistocene layers in Berlin is usually between 20 and 100 m; in the channels from the Elsterian period, however, they can be up to 250 m thick (figure 3).

After the end of the Pleistocene some 12,000 years ago, the Holocene began. Peats, gyttjas, and dune and river sands were deposited, which in some localities had very great thicknesses, often of more than 10 m.

Statistical Base

For the drafting of this Geological Outline, one source consisted of the following older cartographic materials in various scales:

- **Geologische Karten von Berlin** [Geological maps of Berlin] **1:10,000**, (1956 – 1991)
- **Geologische Spezialkarten** [Geological special maps] **1:25,000**, (1878 – 1937)
- **Lithofazieskarte Quartär** [Lithofacies map, Quarternary], **1:50,000**, (1971 – 1988)
- **Geologische Übersichtskarte von Berlin und Brandenburg** [Geological overview map of Berlin and Brandenburg] **1:100,000**, (1995)
- **Geologische Übersichtskarte** [Geological overview map] **1:200,000; sheet CC 3942 – Berlin**, (1998)

The other source was the layer directory of approx. 140,000 drillings as of 2006, from the geological data base of the Berlin State Geological Service, of which excerpts were evaluated.

Methodology

As agreed, the loose rock occurring in the North German Plain is depicted to a depth of 2 m. The material at deeper levels, the geological structure and the deposit conditions and disturbances can be obtained from the profile sections and the explanatory text.

The representation of the various units in the geological outline is carried out for the surface to a **depth of 2 m**, and 5 m in exceptional cases. If various different layers occur in this stratum, e.g. sand over glacial till, only the dominant layer is shown here, for the sake of simplicity; in this case, that would be the till.

Table 2 shows the compiled geological units, distinguished according to stratigraphics (age, chronological sequence), genesis (origin) and petrographics (description of rock).

Tab. 2: Units shown on the geological outline map

Stratigraphics		Genesis		Petrographics (map units)
Quaternary	Holocene		Bog and lake deposits	Peat, gyttja
	Holocene/ Pleistocene		Dune formation	Sand of dunes
	Pleistocene	Weichselian Glaciation	Glacial spillway ancillary valley formation	Valley sand
			Formation of plateaus	Boulder clay, glacial till Meltwater sand
Tertiary	Lower Oligocene		Marine deposits	Clay to silt (rupelton)

Tab. 2: Units shown on the geological outline map

The cover layer (topsoil and deposits) is not shown on the map.

Map Description

Today's surface formations in Berlin were largely shaped by the Weichselian Glaciation, which has also marked the city morphologically (Figure 4).

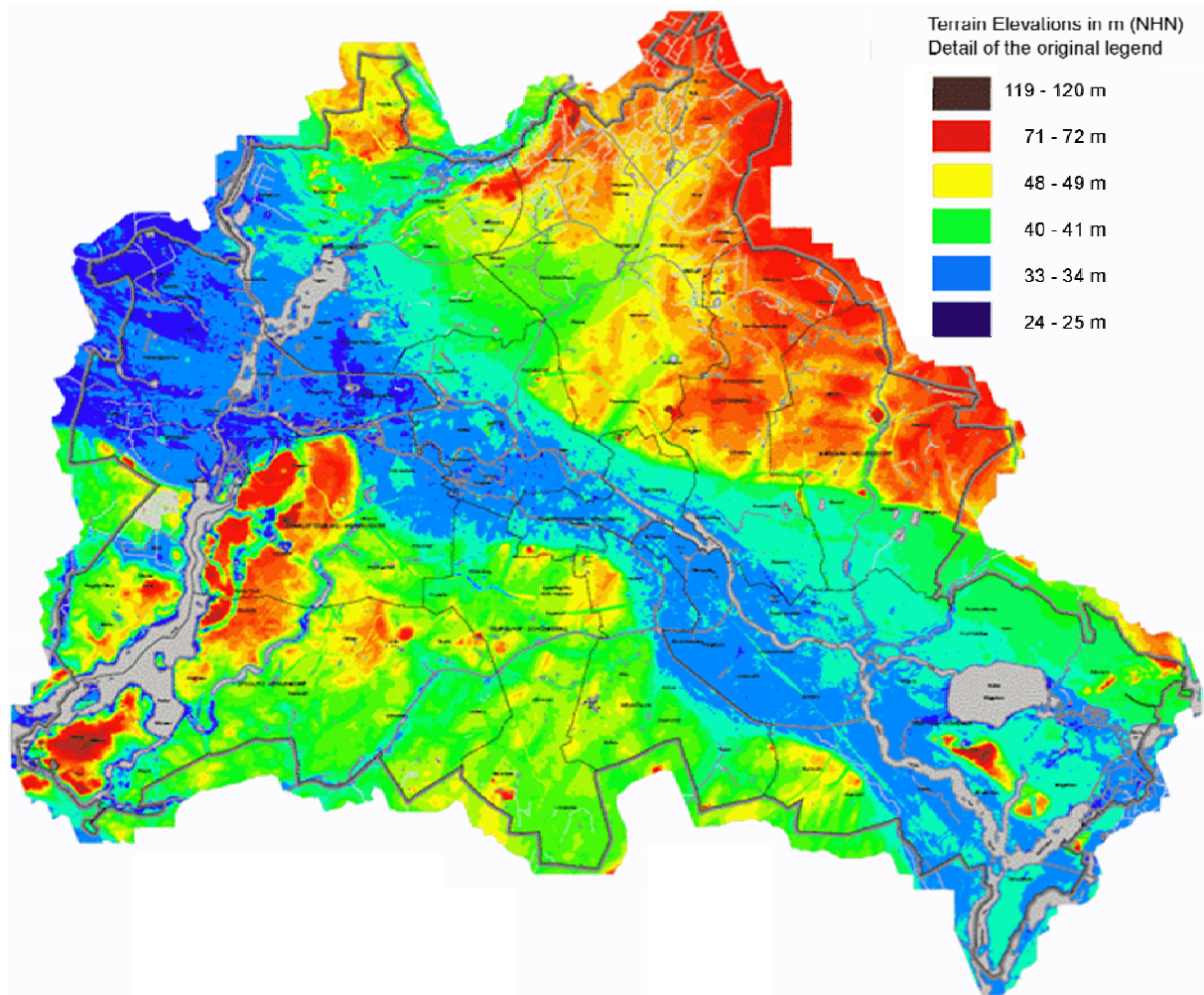


Fig. 4: Morphological outline of Berlin (Environmental Atlas Map 01.08 Terrain Elevation, 2010 edition)

In terms of natural areas, the city appears to be divided into three parts, the **Barnim Plateau** in the north, the **Teltow Plateau**, together with the **Nauen Plate** in the South, and the morphologically lower lying **Warsaw-Berlin Glacial Spillway** (Figure 5). The geomorphological image is supplemented by the low-lying areas of the **Panke Valley** and the **Havel Chain of Lakes**.

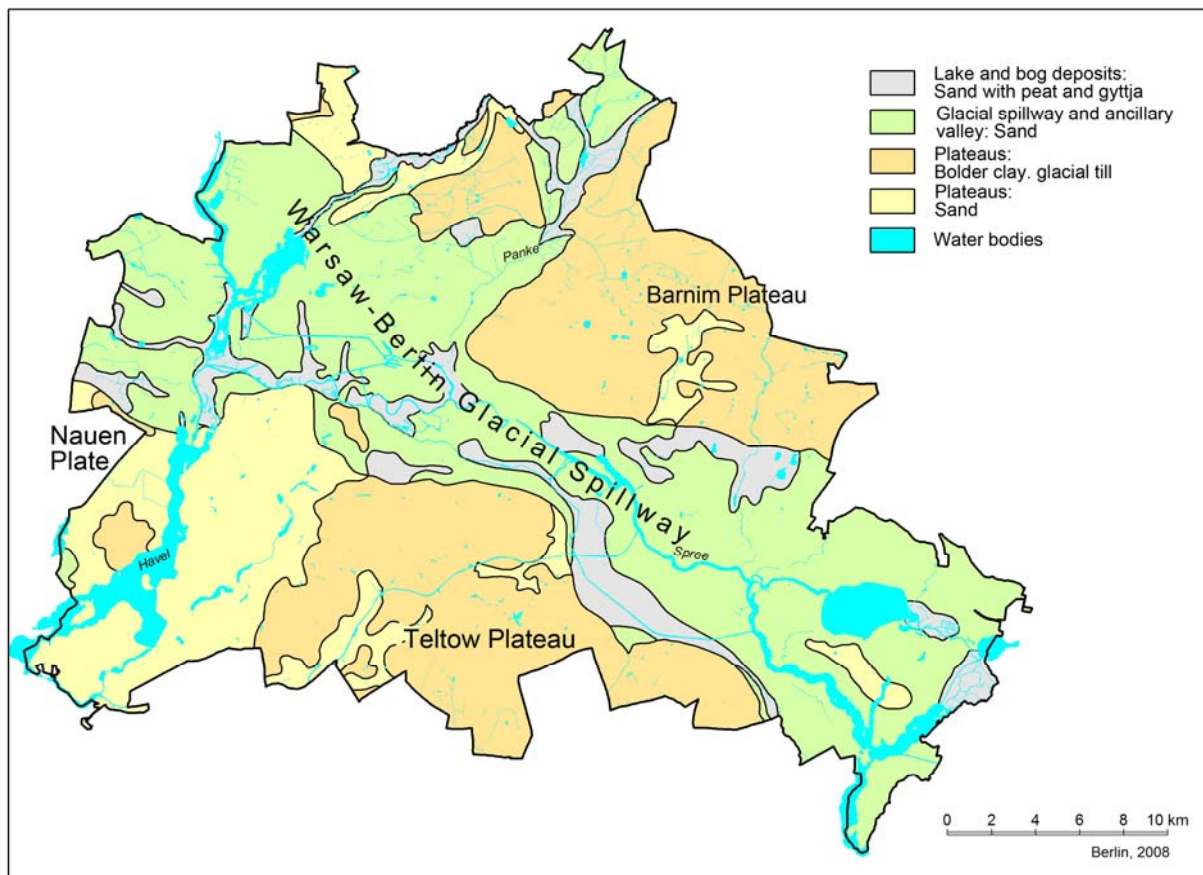


Fig. 5 Natural spatial/geomorphological divisions of Berlin

The **Barnim Plateau**, the ground moraine plate from the Weichselian Glaciation, is located in the northern part of Berlin. Here, the ground moraine and the Weichselian meltwater sands alternate in characterizing the surface. This ground moraine largely directly covers the older ground moraine from the Saalian Glaciation. However, meltwater sands which separate these two ground moraines are also very widespread.

In the Northeast, the Plateau is crossed by the **valley of the Panke**, with its valley sand deposits, which runs roughly north-northeast to south-southwest. In the village of Lübars in the borough of Reinickendorf, there is a small area in which the tertiary rupelton clay is close to the surface, due to salt-tectonic processes; otherwise, this clay occurs only at greater depths of approx. 150 to 200 m below the surface. Starting in the mid-19th century, this clay was quarried in an opencast process, and baked into bricks in a circular kiln built on-site. Such famous buildings as Berlin's Red City Hall and the Reinickendorf Borough Hall were built in part with these bricks. After the quarry was closed, groundwater flowed into the pit, creating a lake which is today used as a swimming facility.

The **Warsaw Berlin Glacial Spillway** is characterized particularly by the accretion of glaciofluvial or fluvial sand and outwash gravel of great thickness – up to 50 m; it passes through the city from east to west, crossing the inner-city area. At some places, saltation load horizons and glacial till residues were deposited (ASSMANN 1957). Locally, these valley sands are covered by younger Holocene sediments (sands, peats and gyttjas mixed with organic material, as well as dunes) of various thicknesses, some very great.

The southern area is occupied by the **Teltow Plateau**, and west of the Havel by the **Nauen Plate**, both of which are also ground moraine plates from the Weichselian Glaciation. In the eastern part of the Teltow Plateau, the ground moraine and meltwater sands from the Weichselian glaciation alternate in characterizing the surface. This ground moraine largely directly covers the older ground moraine from the Saalian Glaciation. However, meltwater sands which separate these two ground moraines are also very widespread. The western part is largely characterized by very thick meltwater sand sequences. The Nauen Plate consists in approximately equal parts of meltwater sands and ground moraine.

The units on the map can be described as follows, starting with the older, lower layers and working up

to the younger, upper ones:

Sediments of the Tertiary

Rupelton

Only in a small area in the village of Lübars east of Hermsdorf Lake can tertiary sediments be found in the form of rupelton clay from the Oligocene.

Petrographically, this is a light grey to olive green, calcium-containing silty clay/clayey silt with a small share of fine sand. The sporadic occurrence of lenses of limestone ("concretions", Ger.: *Septarien*) created by accretive crystallization have given it the name "septarian clay".

Due to its clay content, the rupelton forms a binding sediment with very low water conductive capacity (an aquaclude). In Berlin, with its great thickness of approx. 80 m, it forms the barrier between the lower saltwater and the upper freshwater layers, as it does in large parts of northern Germany.

Sediments of the Quaternary

Meltwater sand

This category of glaciofluvial sediments includes all meltwater sands deposited above, within or below the Weichselian-era ground moraine. They were transported by the water melting off the glacier, and deposited in various forms and grain sizes.

The meltwater sands widely distributed across the ground moraine may contain a share of late Pleistocene to Holocene outwash mass. They may consist of a quite variegated spectrum of grain sizes, ranging from silty fine sand coarse sand, and can also be of variable thickness.

Locally, meltwater sands are widely distributed under the Weichselian ground moraine. Inside this area, glaciofluvial sands of fine to course size separate the upper, Weichselian-era ground moraine from the lower Saalian-era ground moraine.

Moreover, within the glacial till, interglacial sands occur, usually only locally, and generally only of minimal thickness. These sands cannot be distinguished petrographically from the glaciofluvial meltwater sands, so that the delimitation between the Weichselian and the Saalian-era tills is often difficult.

Due to their cohesionless nature, meltwater sands have a great water conductive capacity and are therefore good groundwater conductors.

Boulder clay, glacial till of the Weichselian Glaciation

The moraines of the Brandenburg Phase of the this Weichselian Glaciation, in some places covered by meltwater sands, forms a coherent horizon of glacial till of between five and 10 m in thickness, sometimes more. Petrographically, it consists mostly of weakly clayey, silty, calcium-containing sands with a low share of gravel and stones, and only rarely dropstones. On the surface, the till has been de-calcified by the effects of the weather and has turned into boulder clay.

Due to their share of clay and silt, the till and boulder clay constitute a binding sediment with a low level of water conductive capacity (an aquitard).

In the areas covered with sand and in the edge of the plateaus, the thickness of the ground moraine has generally been reduced by erosion.

The dropstones carried by the glacier or eroded out of the ground moraine can be seen all over the city.

Valley sands in the area of the glacial spillway and the ancillary valleys

As the Weichselian ice continued to melt, the **Warsaw-Berlin Glacial Spillway** was created. Its development occurred in a number of phases. Both the meltwater of the older period of the the Weichselian Glaciation (Frankfurt Phase) and of the younger period (the Pomeranian glacial edge) use the valley courses as drainage channels leading north-westward toward the Elbe and the North Sea.

In the Glacial Spillway, the valley sand is generally more than 10 m thick; below it are the sediments from the Saalian Glaciation. In the upper layers, the valley sand is fine to medium grained, in some cases silty; at greater depths, it becomes coarser, and more often includes gravelly admixtures. Both the Weichselian and the Saalian ground moraines have been largely eroded, so that the sand from the Weichselian Glaciation in most cases rests directly on older sand layers, sometimes as much as 50 m

deep.

In the **Panke Valley**, an ancillary valley of the Glacial Spillway, this same sequence of sand can be seen, although its thickness is seldom greater than 10 m. Locally, glacial till deposits can be seen which are interpreted as erosion residues both of the Weichselian and the Saalian ground moraines.

Due to their cohesionless nature, the valley sand has a great water conductive capacity and is therefore a good groundwater conductor.

Sediments of the Holocene

Dunes

During the late phase of the **Weichselian Glaciation**, which was still cold, dry and poor in vegetation, there was a build-up of dunes and fly-sand fields. Major fly-sand areas and extensive dune fields are no rarity in Berlin, and often occur at the edges of the major lowlands in the Glacial Spillway. In Tegel, Spandau and Friedrichshagen, large and very high dunes (up to 30 m), such as the Ehrenpfortenberg, were raised up by the winds. In Spandau, parabolic dunes were created. The ground moraine areas remained largely free of dunes, since the binding material deposited there was little or not at all suited for fly-sand formation. The dune sands consist of a grain size spectrum ranging from fine to medium sand.

Peat, gyttja

With the melt of the last dead ice of the Weichselian Glaciation, hollow spaces were created in which various types of gyttja and/or humus sand formations arose from the sedimentation which already began during the late Weichselian phase, and continued into the Holocene. In some places, these deposits can be up to 40 m thick, as on Museum Island or at Tiefwerder. The Laach Lake tufa, a deposit of no great thickness resulting from a volcanic eruption in the Eiffel Mountains in 9350 BC, has been attested in such gyttjas at some places in Berlin (PACHUR & RÖPER 1984 and LIMBERG 1991). In connection with the continuing land forming of lakes and other bodies of water, large-scale peat bogs and peat-earth areas (mixture of organic and mineral substance) arose in the lowlands and also on the plateaus.

The Holocene formations are distributed with small or large coherent areas across the entire city, however primarily in the lowlands of the Glacial Spillway and in the valleys on the plateaus. The fluvial floodplain sediments of the Spree can often not be distinguished from the glacial spillway sediments, due to anthropogenic changes in the landscape (construction, the channelling of watercourses).

Anthropogenic changes in the landscape

The landscape of the Berlin metropolitan area has been permanently overformed by human activity. Since the Neolithic, some 5000 years ago, people have been continually impacting upon the landscape. In the Middle Ages, large-scale forest clearance led once again to drifting of sands and to some formation of dunes in the Glacial Spillway (LIMBERG 1991). The damming of water to drive mills and the drainage of swampland during the Middle Ages caused changes in the groundwater situation which affected the development of bogs. Natural and artificial changes in watercourses led to the emergence of dead river arms. Moats were dug around the city, and later filled in again. In the Glacial Spillway, which at that time still consisted largely of swampland, embankments (Ger.: *Damm*) such as the Kurfürstendamm or the Kaiserdamm were built and reinforced to carry roadways and railways. Later, landfill was dumped in low-lying areas. Numerous old gravel quarries in the Glacial Spillway are today artificial lakes or groundwater ponds (Great Spekte Lake, Habermann Lake). Construction rubble, especially that resulting from wartime bombing, was piled up as dumps or hills, such as the Teufelsberg or the hill in the Prenzlauer Berg Public Park. Ultimately, urbanization so overformed large areas of the landscape within the city that in many cases, the original morphological or geological formations can be recognized only with great difficulty.

Major construction projects in the Glacial Spillway which required long-term lowering of the groundwater level have in some areas with very thick Holocene-era organogenic layers (bad construction land) led to settling and major damage to buildings. For example, in the "wet triangle" in the boroughs of Charlottenburg and Mitte (around the old Friedrichstadtpalast), entire houses have had to be torn down, since the long-term lowering of the groundwater caused by underground metro construction has on the one hand led to settling of the terrain, and on the other go damage to the timber piles upon which the buildings were built.

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